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**Ocean Optics Protocols For Satellite Ocean Color Sensor
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Radiometric Measurements and Data Analysis Protocols

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Chapter 3

Above-Water Radiance and Remote Sensing Reflectance Measurement and Analysis Protocols

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3.1 INTRODUCTION

As an alternative to the in-water methods of Volume III, Chapter 2, water-leaving radiance can be measured from the deck of a ship. A shipboard radiometer is used to measure radiance $L_{\text{sfc}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o)$ emanating from the sea surface at zenith angle θ (usually chosen between 30° and 50°) and azimuth angle ϕ (usually chosen between 90° and 180° away the sun's azimuth ϕ_o). In the convention used here, azimuth angles ϕ are measured relative to the sun's azimuth, *i.e.* $\phi_o = 0$.

The surface radiance measured with a radiometer having a solid-angle field of view (FOV) of Ω_{FOV} sr may be expressed, following Mobley (1999), as

$$L_{\text{sfc}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o) = L_{\text{w}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o) + \rho L_{\text{sky}}(\lambda, \theta_{\text{sky}}, \phi_{\text{sky}} \in \Omega'_{\text{FOV}}; \theta_o). \quad (3.1)$$

$L_{\text{w}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o)$ is water-leaving radiance centered at angles (θ, ϕ) and averaged over Ω_{FOV} [as weighted by the radiometer's directional response function (see Volume II, Chapters 2 and 3)]. $L_{\text{sky}}(\lambda, \theta_{\text{sky}}, \phi_{\text{sky}} \in \Omega'_{\text{FOV}}; \theta_o)$ is sky radiance measured with the radiometer looking upward at angles $(\theta_{\text{sky}}, \phi_{\text{sky}})$. In practice, θ and θ_{sky} are numerically equal angles in the nadir and zenith directions, respectively, and the sea and sky viewing azimuths $\phi = \phi_{\text{sky}}$. The reflectance factor ρ is operationally defined as the total skylight actually reflected from the wave-roughened sea surface into direction (θ, ϕ) divided by sky radiance measured with the radiometer from direction $(\theta_{\text{sky}}, \phi_{\text{sky}})$, both quantities being averaged over Ω_{FOV} (Mobley 1999). Remote sensing reflectance is then determined, using water-leaving radiance calculated from (3.1), as

$$R_{\text{RS}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o) = \frac{L_{\text{w}}(\lambda, \theta, \phi \in \Omega_{\text{FOV}}; \theta_o)}{E_{\text{s}}(\lambda; \theta_o)}, \quad (3.2)$$

where $E_{\text{s}}(\lambda; \theta_o)$ is incident spectral irradiance measured above the sea surface. All of the above variables vary with solar zenith angle θ_o .

A simplified notation is used in Volume III, Chapter 2 (and elsewhere in the protocols) when discussing water leaving radiance $L_{\text{w}}(\lambda)$ and remote sensing reflectance $R_{\text{RS}}(\lambda)$ derived from in-water profile measurements of $L_{\text{u}}(z, \lambda)$. Because $L_{\text{u}}(z, \lambda)$ is measured in water viewing the nadir direction, $L_{\text{w}}(\lambda)$ represents radiance leaving the surface in the zenith direction $(\theta, \phi) = (0^\circ, 0^\circ)$. Therefore, $L_{\text{w}}(\lambda)$ in, *e.g.*, Volume III, Chapter 2 corresponds to $L_{\text{w}}(\lambda, 0, 0 \in \Omega_{\text{FOV}}; \theta_o)$, and $R_{\text{RS}}(\lambda)$ to $R_{\text{RS}}(\lambda, 0, 0 \in \Omega_{\text{FOV}}; \theta_o)$, in the present notation.